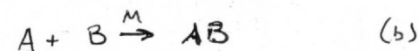


I. Kinetics:



(1) Exothermic ?

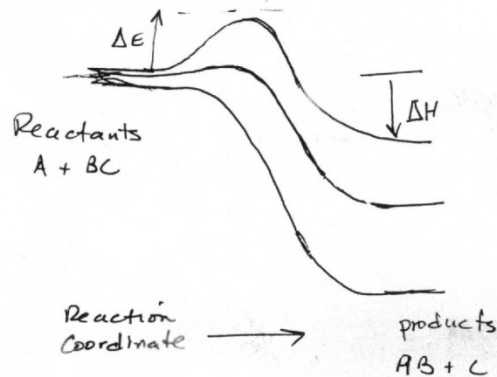
IF $\Delta H > 5 \text{ kcal/mole}$, likely unimportant

(2) Type: · Radical-Radical - (a) or (b)

· Radical-Molecule (a)

· Molecule-Molecule RARE!

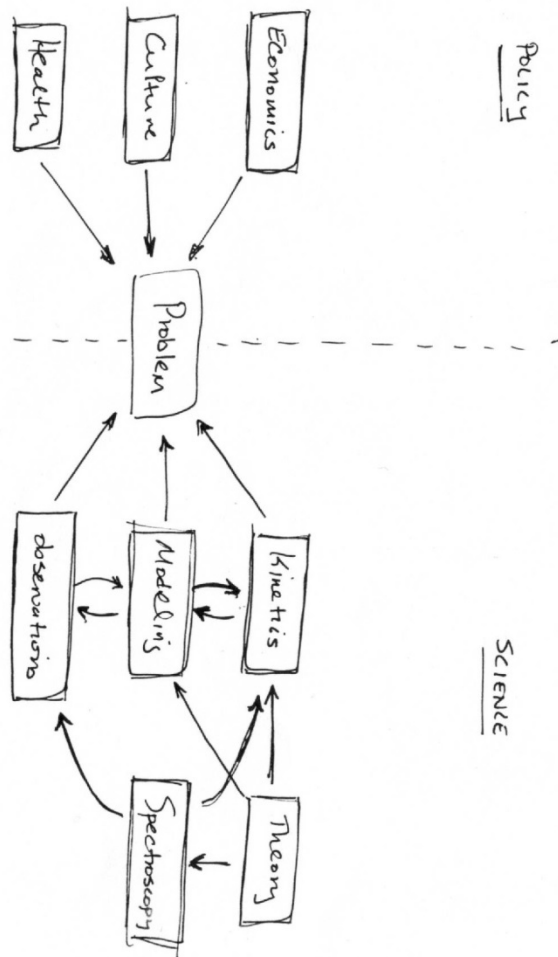
(3) Potential energy surface - bimolecular reaction



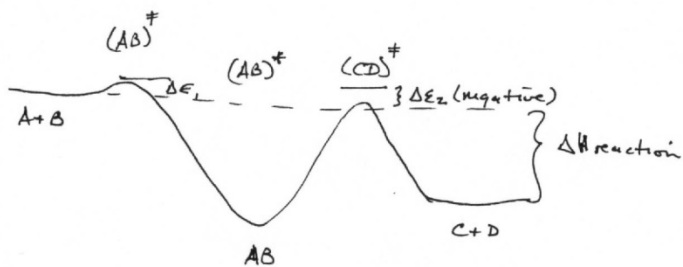
General trend
(not perfect)
decreasing ΔE
for increasing
 ΔH

$$k = A e^{-\Delta E/RT}$$

A = "Preexponential factor"
 ΔE = "Activation Energy"

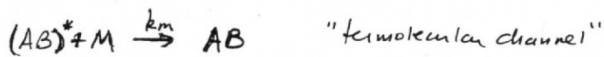
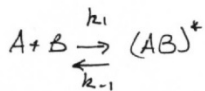


Mixed Behavior:



$$k^{\text{II}} = k_1 \left\{ \frac{k_2 + k_m[M]}{k_{-1} + k_2 + k_m[M]} \right\}$$

where:

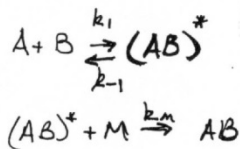
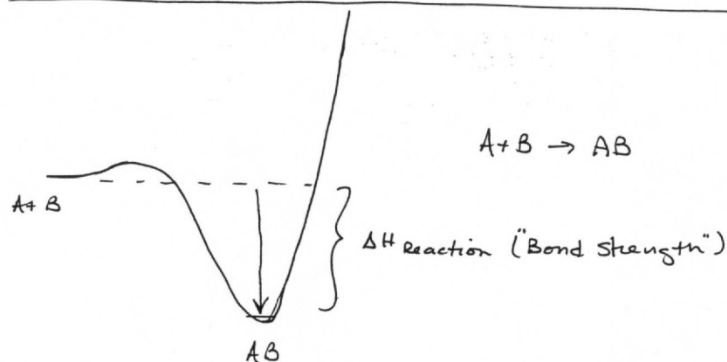


$$k_1 = A_1 e^{-\Delta E_1/RT}$$

$$k_2 = A_2 e^{-\Delta E_2/RT}$$

$k_{-1} \sim T$ -independent (simple bond fission from excited-state complex)

(4) Potential Energy Surface, termolecular Reaction



$$k^{\text{II}} = \frac{k_1 k_m [M]}{k_{-1} + k_m [M]} \quad \text{from steady-state } (AB)^*$$

where

$$\frac{d[CA]}{dt} = -k^{\text{II}} [A][B] \left(= -\frac{d[AB]}{dt} \right)$$

low P: $k^{\text{II}} \approx \frac{k_1 k_m [M]}{k_{-1}}$ linear in [M]

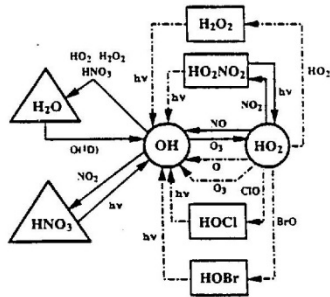
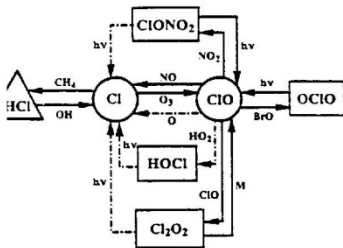
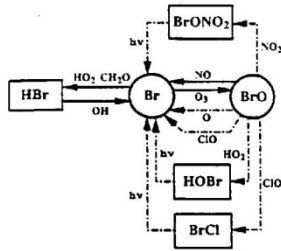
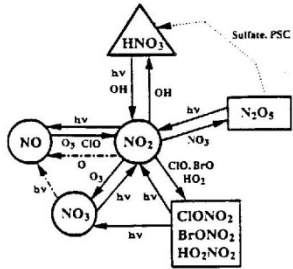
high P: $k^{\text{II}} \approx k_1$ (high pressure limit)

(6) Photolysis

Photochemistry: Stratosphere

$$J(h, T, \phi) = \int_0^{\infty} F_{sun}(\lambda, h, \phi) \cdot \sigma(\lambda, T) \cdot \phi(\lambda, T, p) d\lambda$$

= photolysis rate (function of Temperature, altitude, solar zenith angle and pressure).



F • Solar flux is dependent on the ozone column between point of interest and Sun, so it depends on wavelength, altitude, and solar zenith angle.

σ • Cross section is a molecular property that depends on Temperature and wavelength (mainly)

φ • Quantum yield depends on wavelength (mainly) and to a much lesser extent on Temperature & pressure.

It is easy to estimate a photolysis rate with tabulated values of solar flux.

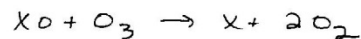
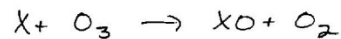
It is hard to calculate an accurate photolysis rate (x2) w/o a model that includes Rayleigh scattering, multiple scattering, ozone distribution, etc.

- 21 photolysis rates
- 33 reactions
- 28 species

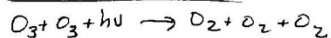
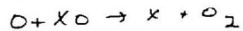
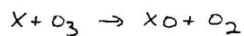
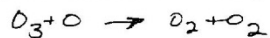
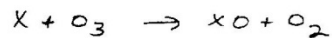
This will get you 90% of the result for most of Stratosphere.

△ = long-lived reservoir
○ = radical

Ozone destruction

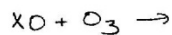
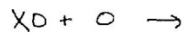


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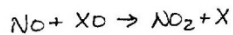


NOTE:

Typically,



all RATE-DETERMINING
STEPS because

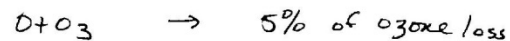
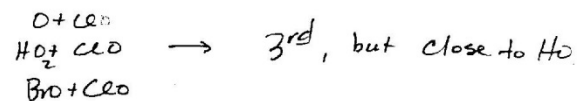
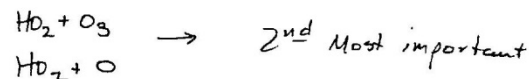
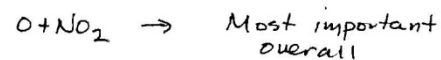


IS FAST!

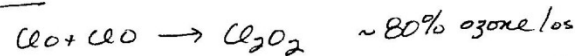
OZONE LOSS:

$$\begin{aligned} \text{Loss} = & 2k[O](ClO) + 2k[ClO](HO_2) + 2k[ClO](BrO) \\ & + 2kClO + 2k[HO_2]^2 + 2k[HO_2](BrO) \\ & + 2k[HO_2](O) + 2k[HO_2](O_3) + 2k[BrO](O) \\ & + 2k[NO_2](O) + 2k[O](O_3) \end{aligned}$$

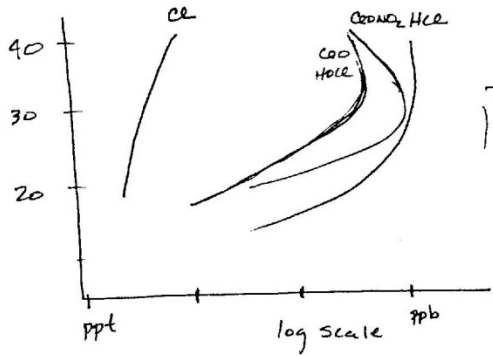
Typically,



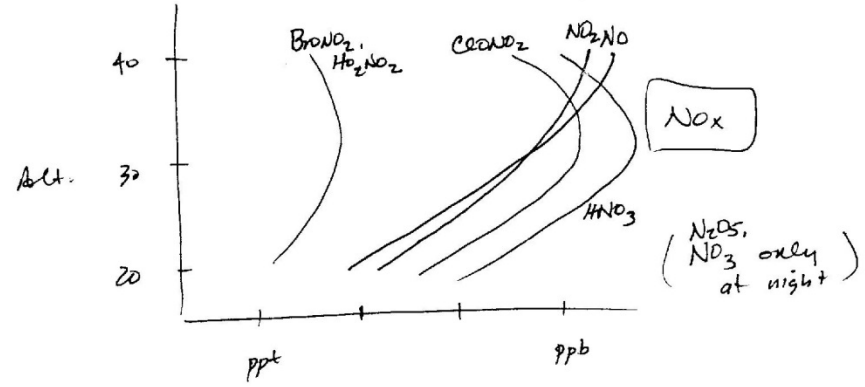
Antarctica:



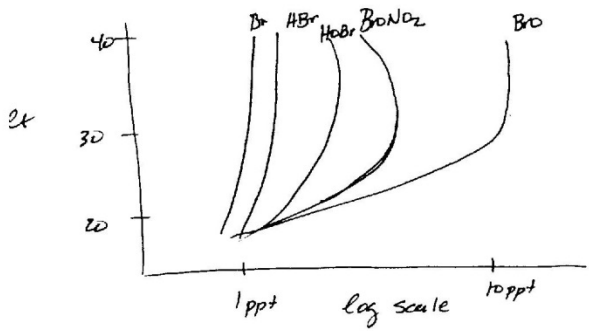
OBSERVATIONS:



ClO₂ important
 Cl
 polar regions
 in winter



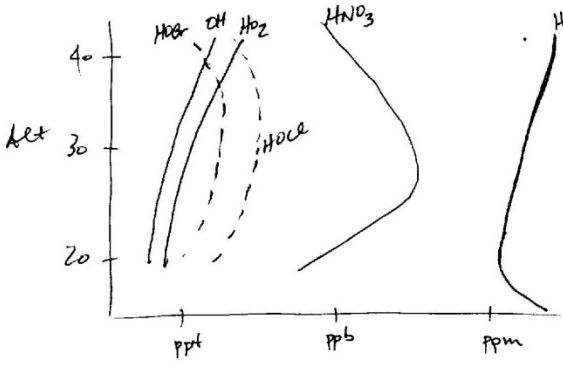
(NO₂,
 NO₃ only
 at night)



BrO important
 in polar
 regions in
 winter.

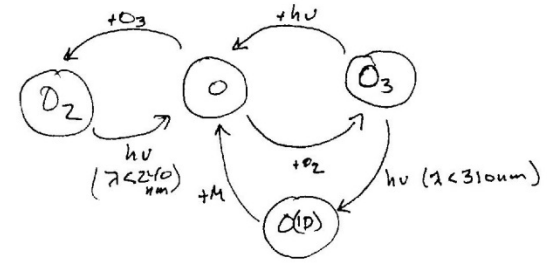
STEADY-STATE, (EXAMPLE)

$$\frac{[BrO]}{[Br]} = \frac{k[O_3]}{k[ClO] + k[O] + k[NO]} \quad (k\text{'s all different})$$



HO₂NO₂?
 H₂O₂?

CHAPMAN:



Heterogeneous Chemistry

(Page 203)

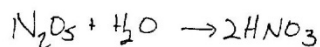
$$k = \frac{\gamma}{4} \left(\frac{8kT}{\pi m_x} \right)^{1/2} \text{S.A.}$$

"Sticking coefficient" (loosely) \rightarrow γ

"Molecular speed" \rightarrow $\left(\frac{8kT}{\pi m_x} \right)^{1/2}$

atmosol surface area. \rightarrow S.A.

EXAMPLE:



$$\gamma = 0.1$$

$$M_{\text{N}_2\text{O}_5} = 108 \text{ AMU. } (= 1.79 \times 10^{-25} \text{ kg/molecule})$$

$$K = 1.381 \times 10^{-23} \text{ J/K}$$

$$\text{at } 220 \text{ K, } \left(\frac{8kT}{\pi m_x} \right)^{1/2} = 208 \text{ m/s}$$

$$k = \frac{0.1}{4} (208 \text{ m/s}) \cdot \text{S.A.}$$

$$\text{typical surface area } \sim 1 \mu\text{m}^2/\text{cm}^3$$

~~etc.~~

So,

$$k = \frac{0.1}{4} (208 \text{ m/s}) \cdot 1 \mu\text{m}^2/\text{cm}^3$$

$$k = 5.2 \times 10^{-6} \text{ s}^{-1}$$

(lifetime of $\text{N}_2\text{O}_5 \sim 2 \text{ days}$)